



Sum of Hourly vs. Daily Penman-Monteith Grass-Reference Evapotranspiration under Semi-arid and Arid Climate

Koffi Djaman^{1*}, Koudahe K², Lombard K¹ and O'Neill M¹

¹Department of Plant and Environmental Sciences, Agricultural Science Center at Farmington, New Mexico State University, Farmington, USA

²ADA Consulting Africa, Lomé, Togo

Abstract

Under changing climate and the global warming, accurate estimation of crop evapotranspiration might help improving water management and maintaining or improving water productivity. The objective of this study was to compare the sum of hourly ETo with the daily average ETo using the standardized ASCE Penman-Monteith equation under semi-arid dry climate. Pair comparison was performed between the sum of hourly reference evapotranspiration (ETo) and the daily average ETo based on the data from four automated weather stations in the State of New Mexico (USA) for the period of 2009-2017. The results showed good agreement between the sum of hourly ETo and daily average ETo with the regression slopes varying from 1.00 to 1.10 and coefficients of determination from 0.63 to 0.97. The least difference between the hourly and daily results was found at Leyendecker. The RMSE values were 0.98, 0.64, 0.21 and 0.35 mm/day at Fabian Garcia, Farmington, Leyendecker and Tucumcari, respectively. The adoption of the hourly ETo estimation method leads to 16.6, 5.5, -0.2, and 4.1% higher annual ETo at Fabian Garcia, Farmington, Leyendecker and Tucumcari respectively, and can help on accurate evapotranspiration estimation to meet crop water requirement under sustainable agriculture.

Keywords: Evapotranspiration; Sum of hourly; Penman-Monteith; Semi-arid climate

Introduction

Crop evapotranspiration is the main water losses from the hydrological cycle and is a very important parameter considered for water planning and management under the agricultural and environmental studies. Under global warming the world temperature is rising and the fresh water for crop and fiber production is diminishing [1], agriculture may suffer from drought due to the decreasing trend in annual precipitation with inter-annual variability in most parts of the world [2-9] while abundant precipitation might occur in other parts of the world with increased extreme precipitation events [6,10-13]. The South western United States characterized by the semi-arid and arid climate, suffers from drought with increasing temperature, decreasing water availability, high evapotranspiration and percolation losses, and increase in forest woodland mortality [14-16]. In the State of New Mexico particularly, with the ongoing ground water depletion, decreasing trend in precipitation, decreasing stream flow [17], and increasing population, a smart agriculture should be of first choice to cope with the changing climate and maintain and or improve water productivity. Ahadi et al. [18] reported irrigation efficiency varying from 11 to 95% across the lower Rio Grande Basin in New Mexico and that averaged 64%. More effort should be done by irrigation managers to much crop water use to the applied irrigation. Accurate crop evapotranspiration estimation is therefore critical for water management under this arid environment. As most of the irrigation scheduling programs are based on reference evapotranspiration (ETo) and with the tendency of drip irrigation adoption for minimizing water loss and increasing irrigation efficiency, there is a need to investigate the accuracy of using daily average weather variables against the hourly weather variables. Numerous ETo estimation models have been developed in addition to the direct measurement methods [19] and the Penman-Monteith method is worldwide recommended for its accuracy under different climatic conditions [20,21].

Some studies reported uncertainty when applying the Penman Monteith equation using daily, weekly or monthly weather data [22,23], while the distribution of climatic variable during the day time and the night time can be source of error in estimating daily ETo [24]

because the nighttime ETo can be non-negligible and equivalent to as high as 30% of the daily ETo [25,26]. Irmak et al. [24] reported 5 to 8% ETo overestimation when using average daily weather data compared to the hourly data from Bushland (Texas), North Platte (Nebraska), Santa Rosa, and Twitchell Island (California) and 3% underestimation at Santa Barbara (California). Djaman et al. [27] reported ETo overestimation range from 4.9 to 8.3% across the semi-arid region in Senegal. In contrast, Bakhtiari et al. [28] found ETo overestimation from 7.4 to 47.6% in Iran when using hourly climate variables data. From the non-consistency of the equivalence of the daily ETo estimates using the average daily data and hourly data, it urges to investigate the accuracy of these time steps calculations under the semi-arid dry and arid condition where water resources are the main limiting factor for crop production. The objective of this study was to compare the daily ETo estimates using average daily and the hourly climate variables data for the semi-arid dry and arid climatic condition in New Mexico.

Materials and Methods

This study was conducted at four weather stations across New Mexico (USA) at Fabian-Garcia, Farmington, Leyendecker and Tucumcari for the period of 2009-2017. The geographical coordinates and the long term average climatic variables are summarized in Table 1. Minimum temperature (Tmin), maximum temperature (Tmax), average temperature (Tmean), minimum relative humidity (RHmin), maximum relative humidity (RHmax), average relative humidity (RHmean), average wind speed (u_2), and solar radiation (Rs) were collected on the hourly basis and averaged over daily time step from

***Corresponding author:** Koffi Djaman, Department of Plant and Environmental Sciences, Agricultural Science Center at Farmington, New Mexico State University, Farmington, USA, Tel: +1505 960 7757; E-mail: kdjaman@nmsu.edu

Received January 28, 2018; **Accepted** February 05, 2018; **Published** February 13, 2018

Citation: Djaman K, Koudahe K, Lombard K, O'Neill M (2018) Sum of Hourly vs. Daily Penman-Monteith Grass-Reference Evapotranspiration under Semi-arid and Arid Climate. Irrigat Drainage Sys Eng 7: 202. doi: [10.4172/2168-9768.1000202](https://doi.org/10.4172/2168-9768.1000202)

Copyright: © 2018 Djaman K, et al. This is an open-access article distributed under the terms of the Creative Commons Attribution License, which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.

| Parameters | Fabian Garcia | Farmington | Leyendecker | Tucumcari |
|--------------------------|---------------|------------|-------------|-----------|
| Latitude (Deg. North) | 32.28 | 36.69 | 32.20 | 35.20 |
| Longitude (Deg. West) | -106.77 | -108.31 | -106.74 | -103.69 |
| Elevation (m) | 1186 | 1720 | 1176 | 1246 |
| Wind speed (m/s) | 1.82 | 2.45 | 1.81 | 3.30 |
| Tmax (°C) | 26.50 | 20.94 | 26.12 | 24.07 |
| Tmin (°C) | 9.82 | 4.33 | 7.67 | 7.52 |
| Tmean (°C) | 18.16 | 12.63 | 16.89 | 15.79 |
| Rhmax (%) | 65.21 | 68.63 | 77.34 | 71.78 |
| Rhmin (%) | 18.60 | 19.70 | 19.09 | 22.29 |
| Rhmean (%) | 41.90 | 44.17 | 48.21 | 47.03 |
| Rs (MJ m ⁻²) | 21.23 | 19.84 | 19.63 | 19.80 |

Tmin: Minimum temperature; Tmax: Maximum temperature; Tmean: Average Temperature; Rhmin: Minimum Relative Humidity; Rhmax: Maximum Relative Humidity; Rhmean: Average Relative Humidity and Rs: Solar Radiation.

Table 1: Weather stations with long term average climatic condition.

automated weather stations installed by the New Mexico Climate Center.

Penman-Monteith reference evapotranspiration model: (ASCE-EWRI, 2005)

The hourly ETo values calculation approach was developed by Snyder and Eching [29], following the description in the ASCE-EWRI report [21]. Required data for ETo calculations using ASCE-PM approach included data about site characteristics (latitude and elevation) and weather data, as hourly solar radiation (Rs) (MJ m⁻² h⁻¹), mean air temperature (T) (°C), mean wind speed (u₂) (m s⁻¹) and mean dew point temperature (Td) (°C). Td was calculated as recommended by Allen et al. [20]. Actual vapor pressure (ea (kPa)) was calculated from mean saturation vapor pressure (es) (kPa) and RH (%). Daily grass-reference ET (ETo) was computed using the standardized ASCE form of the Penman-Monteith (ASCE-PM) equation [21]. The Penman-Monteith reference evapotranspiration equation with fixed stomatal resistance values for grass surface is:

$$ETo = \frac{0.408\Delta(Rn - G) + \gamma Cnu2 / (T + 273)(es - ea)}{\Delta + \gamma(1 + Cdu2)} \quad (1)$$

Where: ETo is the reference evapotranspiration (mm/day);

Δ=The slope of saturation vapor pressure versus air temperature curve (kPa°C⁻¹);

Rn=Net radiation at the crop surface (MJ m⁻² d⁻¹);

G=Soil heat flux density at the soil surface (MJ m⁻² d⁻¹);

T=Mean daily air temperature at 1.5-2.5 m height (°C);

u₂=Mean daily wind speed at 2 m height (m s⁻¹);

es=The saturation vapor pressure (kPa);

ea=The actual vapor pressure (kPa);

es-ea=Saturation vapor pressure deficit (kPa);

γ=Psychrometric constant (kPa °C⁻¹);

Cn=Numerator constant that changes with reference surface and calculation time step (900°C mm s³ Mg⁻¹ d⁻¹ for 24 h time steps, and 37°C mm s³ Mg⁻¹ h⁻¹ for hourly time steps for the grass-reference surface);

Cd=denominator constant that changes with reference surface and calculation time step (0.34 s m⁻¹ for 24 h time steps, 0.24 s m⁻¹ for

hourly time steps during daytime, and 0.96 s m⁻¹ for hourly nighttime for the grass-reference surface).

All parameters necessary for computing ETo were computed according to the procedure developed in FAO-56 by Allen et al. [20].

Evaluation criteria

Comparisons were made using the *t*-test, graphics and simple linear regression. A paired sample *t*-test (two-sample for means) was performed to identify any significant difference between the sum of hourly ETo and the daily average ETo estimates at 5% significance level. The null hypothesis was that the daily average ETo estimates and the sum of hourly ETo estimates came from the same population and that mean difference between ETo estimates was zero. The linear regressions were forced through the origin because ideally all equations should produce zero ETo when there is no evapotranspiration. The regression slopes, coefficient of determination (R²), Root mean squared error (RMSE), Percent error (PE), mean bias error (MBE), and absolute mean error (MAE) were also used and the RMSE, PE were calculated as follow:

$$RMSE = \sqrt{\frac{\sum_{i=0}^n (Daily ETo_i - Sum of hourly ETo_i)^2}{n}} \quad (2)$$

$$PE = \frac{RMSE}{Sum of hourly ETo mean} \times 100 \quad (3)$$

$$MBE = n^{-1} \sum_{i=1}^n (Daily ETo_i - Sum of hourly ETo_i) \quad (4)$$

$$MAE = n^{-1} \sum_{i=1}^n |Daily ETo_i - Sum of hourly ETo_i| \quad (5)$$

Where:

daily ETo=ETo calculated at the daily time step;

Sum of hourly ETo=sum of 24 hour ETo calculated at an hourly time step and considered as the standard or measured ETo at the *ith* data point and *n* is the total number of data points.

Results and Discussion

Maximum hourly wind speed, temperature, relative humidity and solar radiation were 11.2 m/s, 41.7°C, 89.7% and 4.22 MJ/m², respectively, at Fabian Garcia, 14.7 m/s, 37.7°C, 100%, and 3.92 MJ/m², respectively, at Farmington, 11.3 m/s 40.9°C, 97.7% and 3.95 MJ/m², respectively, at Leyendecker, and 15.3 m/s, 41.9°C, 97.7% and 4.09 MJ/m² at Tucumcari. Daily average values were 5.3 m/s, 34.2°C, 89.6%, 32.3 MJ/m² for the hourly wind speed, temperature, relative humidity and solar radiation, respectively, at Fabian Garcia, 8.2 m/s, 29.7°C, 96.3% and 33.20 MJ/m² at Farmington, 6.1 m/s, 31.2°C, 87.6% and 32.31 MJ/m² for the respective variables at Leyendecker, and 9.7 m/s, 33.0°C, 97.3% and 33.9 MJ/m² for the respective variables at Tucumcari. It can be drawn that the daily average approach does not account for the hourly abrupt changes in the climate variables.

The sum of hourly ETo showed good agreement with the daily average ETo (Figures 1-3). The regression slope between the sum of hourly ETo and the daily average ETo and the coefficient of determination were 1.104 and 0.63 at Fabian Garcia, 1.091 and 0.92 at Farmington, 1.002 and 0.97 at Leyendecker and 1.028 and 0.97 at Tucumcari. Overall, error related to the ETo estimates using daily average weather data is 10, 9, 0.2 and 3% at Fabian Garcia, Farmington, Leyendecker and Tucumcari, respectively. The data showed ETo slight underestimation during the year at Fabian Garcia (Figures 1 and 2), and from April to August in Farmington (Figure 2). However, the peak

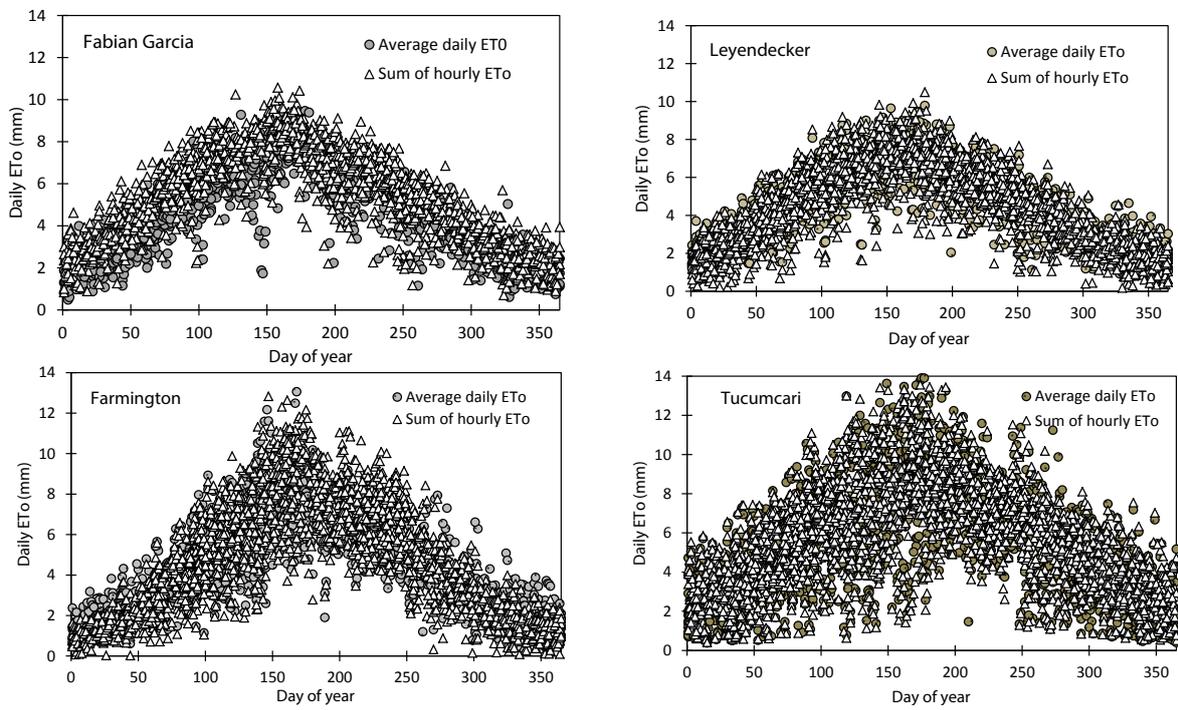


Figure 1: Comparison of the sum of hourly ETo and the average daily ETo at Fabian Garcia, Farmington, Leyendecker and Tucumcari weather stations.

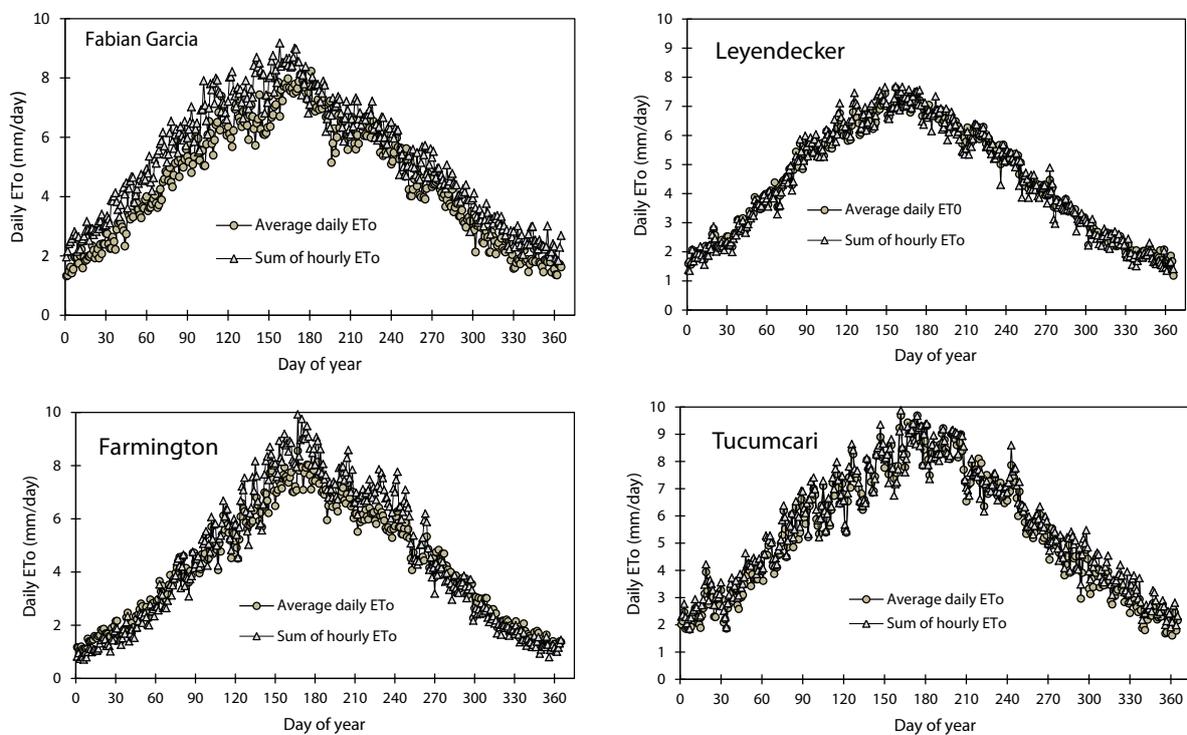


Figure 2: Comparison of the mean of the sum of hourly ETo and the mean of the average daily ETo at Fabian Garcia, Farmington, Leyendecker and Tucumcari weather stations.

ET_o month is June at all four locations and the sum of hourly showed higher peak ET_o values of 9.2 and 9.3 mm/day at Fabian Garcia and Farmington, respectively, while the daily average ET_o peak values were 8.0 and 8.5 mm/day at the respective locations. The annual mean ET_o using the hourly and the daily average climatic variables was 1938.8 and 1662.1 mm at Fabian Garcia, 1647.0 and 1561.3 mm at Farmington, 1626.4 and 1629.9 mm at Leyendecker and 2043.9 and 1960 mm at Tucumcari. Using hourly data leads to 16.6, 5.5, -0.2, and 4.1% higher annual ET_o compared to the daily average approach. The t-test showed significant difference between the annual mean daily average ET_o and the annual mean sum of hourly ET_o at Fabien Garcia, Farminton and Tucumcari (Table 2). No significant difference between the annual mean daily average ET_o and the annual mean sum of hourly ET_o at Legendecker (Table 2). For practical irrigation water requirement estimation and irrigation management the hourly data should be preferred to the daily average mean to avoid putting crop under water stress that will impact the yield if the data is available at Farmington. The RMSE values were 0.98, 0.64, 0.21 and 0.35 mm/day and were equivalent to PE values of 18.5, 14.2, 4.8 and 6.3% at Fabian Garcia, Farmington, Leyendecker and Tucumcari, respectively (Table 3). The lowest MBE and MAE were obtained at Leyendecker and the greatest

values were observed at Fabian Garcia (Table 3). The application of the hourly ET_o approach considers the diurnal abrupt changes in the climate variables and should be of first choice up to the availability of the data [24-30]. However, the Hourly and daily ET_o estimation approach could be interchangeable at Leyendecker, Tucumcari and Farmington due to the small differences between daily ET_o values revealed by this study at these weather stations.

Djaman et al. [27] reported good agreement between the daily and hourly time steps ET_o with regression slope from 1.02 to 1.08 and R² greater than 0.87. The daily time step ET_o showed ET_o overestimation from 1.3 to 8% as compared to the sum of hourly ET_o across the semiarid and humid region of West Africa [27]. Similar results were reported in the Southern Spain by Gavilan et al. [30] and in the state of Georgia (USA) by Suleiman and Hoogenboom [31]. The hourly ET_o over 24 hours was 5.8 to 47.6% higher than daily ET_o using average daily weather data in Iran [28]. Jia et al. [32] also found that the daily showed good agreement with the sum of 24-hour reference ET_o under the sub-humid continental climate in North Dakotas (USA). Allen et al. [33] indicated that estimating ET_o at hour time step improves the accuracy of ET_o estimation under very variable and dynamics environment in locations with large variation in wind speed and cloud

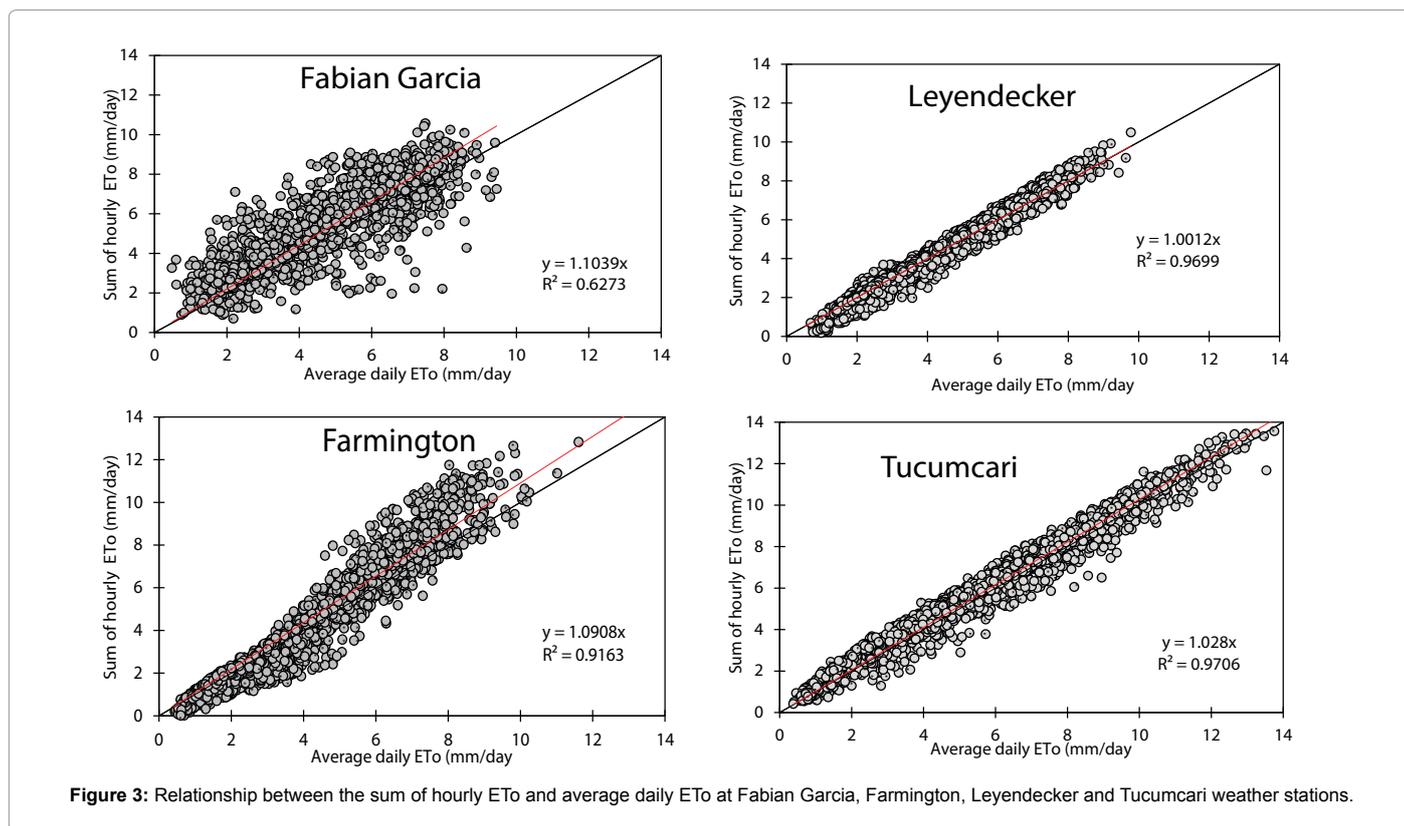


Figure 3: Relationship between the sum of hourly ET_o and average daily ET_o at Fabian Garcia, Farmington, Leyendecker and Tucumcari weather stations.

| Weather station | Mean ET _o | | Variance ET _o | | t-test (one tail) | | P-value | Significance |
|-----------------|----------------------|---------------|--------------------------|---------------|-------------------|------------|---------|--------------|
| | Daily Average | Sum of hourly | Daily Average | Sum of hourly | t-computed | t-critical | | |
| Fabien Garcia | 4.55 | 5.31 | 3.84 | 3.98 | -23.27 | 1.649 | 3E-74 | * |
| Farmington | 4.28 | 4.51 | 4.37 | 6.70 | -7.53 | 1.649 | 2E-13 | * |
| Leyendecker | 4.46 | 4.45 | 3.34 | 3.63 | 0.91 | 1.649 | 0.181 | n.s. |
| Tucumcari | 5.37 | 5.60 | 4.98 | 4.52 | -16.20 | 1.649 | 4E-45 | * |

*Significant at the 5% significance level.

n.s.=non-significant at the 5% significance level.

Table 2: Summary of the paired sample t-test (two-sample for means) statistics for the daily average ET_o versus the sum of hourly ET_o (α=0.05).

| Indices | Fabian Garcia | Farmington | Leyendecker | Tucumcari |
|------------------|---------------|------------|-------------|-----------|
| Regression slope | 1.1 | 1.09 | 1.001 | 1.03 |
| R ² | 0.63 | 0.92 | 0.97 | 0.97 |
| RMSE (mm/day) | 0.98 | 0.64 | 0.21 | 0.35 |
| PE (%) | 18.5 | 14.2 | 4.8 | 6.3 |
| MBE (mm/day) | 0.76 | 0.23 | -0.01 | 0.29 |
| MAE (mm/day) | 0.82 | 0.5 | 0.16 | 0.23 |

R²: Coefficient of Determination; RMSE: Root Mean Squared Error; PE: Percent Error; MBE: Mean Bias Error and MAE: Absolute Mean Error.

Table 3: Statistics summary of the comparison between the sum of hourly and daily Penman-Monteith ETo.

cover. Perera et al. [34] found the best agreement between the hourly and daily results for the FAO-Penman-Monteith version in temperate climates and the ASCE-Penman-Monteith version in the tropical and arid climates while Berengena and Gavilan [35] reported daily ETo underestimation of 3% and 2% by the FAO-Penman-Monteith version ASCE-Penman-Monteith versions in Cordoba (Spain), respectively. Irmak et al. [24] indicated that differences between daily average ETo and sum of hourly ETo were shown under strong, dry, hot winds with advective increases in hourly ETo while the daily average ETo did not account for abrupt changes in wind speed, air temperature, and vapor pressure deficit during the day. Moreover, the night time ETo can be equivalent to 15% of the daily ETo in semiarid and arid environment as reported by Tolk et al. [25]. Caird et al. [26] reported the nighttime transpiration rates of 5-30% of the daily ETo. Seasonal variability in the ratio of the sum of hourly ETo and daily ETo at Fabian Garcia and Farmington is in agreement with the results of Perera et al. [34] in the tropical and arid climates in Australia.

The finding of this study might be useful to crop growers, irrigation managers and environmentalists for water management and planning to optimized crop actual water vs. potential irrigation requirement across the State of New Mexico. With the development of irrigation management technology, producers are encouraged to go for high irrigation frequency especially under the coarse sandy soils with small water holding capacity and the extremely high air temperature during the crop growth and development period similar to New Mexico for effective water and nutrient management [36,37]. The hourly actual evapotranspiration base irrigation might be more efficient due to the slight variability of the surface resistance, and crop coefficient during daytime [38] and the hourly ETo might be much more useful for crop irrigation water requirement for a short period, growth stage or the growing period [24,27,39]. Locally developed hourly and daily crop coefficients data [36,40,41] could be used to improve water management at field, scheme and watershed levels. ETo underestimation might conduct to crop under irrigation that might reduce crop yield and its overestimation will lead to crop fading over irrigation with negative impact on crop yield and promote nutrient leaching.

Conclusion

Comparison between the sum of hourly reference evapotranspiration and the daily average evapotranspiration was performed at four automated weather stations in the State of New Mexico (USA) for the period of 2009-2017 using the standardized Penman-Monteith reference evapotranspiration equation. The results showed good agreement between the sum of hourly ETo and daily average ETo with the regression slopes varying from 1.00 to 1.10, the coefficient of determination from 0.63 to 0.97 and the RMSE values varied from 0.21 to 0.98 mm/day. ETo estimation at hourly basis that accounts for the abrupt changes in the weather variables, improves ETo estimation and might help estimating accurate crop water requirement

and improving water management and water productivity under semiarid dry climate in the State of New Mexico and the neighboring States under similar climatic conditions with water resources as the most limiting factor for food and fiber production.

References

1. Intergovernmental Panel on Climate Change (2015) Climate change 2014: mitigation of climate change. Cambridge University Press 3.
2. Zhai P, Zhang X, Wan H, Pan X (2005) Trends in total precipitation and frequency of daily precipitation extremes over China. *Journal of Climate* 18: 1096-1108.
3. Paredes D, Trigo RM, Garcia-Herrera R, Trigo IF (2006) Understanding precipitation changes in Iberia in early spring: weather typing and storm-tracking approaches. *Journal of Hydrometeorology* 7: 101-113.
4. Nicholson SE (2013) The West African Sahel: A review of recent studies on the rainfall regime and its interannual variability. *ISRN Meteorology*.
5. Sayemuzzaman M, Jha MK (2014) Seasonal and annual precipitation time series trend analysis in North Carolina, United States. *Atmospheric Research* 137: 183-194.
6. Maidment RI, Allan RP, Black E (2015) Recent observed and simulated changes in precipitation over Africa. *Geophys Res Lett* 42: 8155-8164.
7. Rahmat SN, Jayasuriya N, Bhuiyan MA (2015) Precipitation trends in Victoria, Australia. *J Water Clim Chang* 6: 278-287.
8. Fleig AK, Tallaksen LM, James P, Hisdal H, Stahl K (2015) Attribution of European precipitation and temperature trends to changes in synoptic circulation. *Hydrol Earth Syst Sci* 19: 3093-3107.
9. Yu X, Zhao G, Zhao W, Yan T, Yuan X (2017) Analysis of Precipitation and Drought Data in Hexi Corridor, Northwest China. *Hydrology* 4: 29.
10. Wang Y, Zhou L (2005) Observed trends in extreme precipitation events in China during 1961-2001 and the associated changes in large-scale circulation. *Geophys Res Lett* 32.
11. Knapp AK, Hoover DL, Wilcox KR, Avolio ML, Koerner SE, et al. (2015) Characterizing differences in precipitation regimes of extreme wet and dry years: implications for climate change experiments. *Global change biology* 21: 2624-2633.
12. Fischer EM, Knutti R (2016) Observed heavy precipitation increase confirms theory and early models. *Nat Climate Change* 6: 986-991.
13. Rajczak J, Schar C (2017) Projections of Future Precipitation Extremes Over Europe: A Multimodel Assessment of Climate Simulations. *Journal of Geophysical Research: Atmospheres* 122: 10,773-10,800.
14. Garcia LA, Schmitz A, Puente A (2004) Agricultural production trends and the future of the trans-boundary. Rio Grande/Rio Bravo Basin Conference Proceedings. San Antonio, Texas: Woodrow Wilson International Center for Scholars Mexico Institute.
15. Mac Donald GM (2010) Climate change and water in Southwestern North America special feature: Water, climate change, and sustainability in the southwest. *Proc Natl Acad Sci USA* 107: 21256-21262.
16. Harpold A, Brooks P, Rajagopal S, Heidbuchel I, Jardine A, et al. (2012) Changes in snowpack accumulation and ablation in the intermountain west. *Water Resour Res* 48: 11501.
17. Hurd BH, Coonrod J (2008) Climate change and its implications for New Mexico's water resources and economic opportunities. Technical Report 45. Las Cruces, New Mexico: New Mexico State University.
18. Ahadia R, Samani Z, Skaggs R (2013) Evaluating on farm irrigation efficiency across the watershed: A case study of New Mexico's Lower Rio Grande Basin. *Agric Water Manage* 124: 52-57.
19. Djaman K, Irmak S, Futakuchi K (2017a) Daily reference evapotranspiration estimation under limited data in Eastern Africa. *Journal of Irrigation and Drainage Engineering* 143.
20. Allen RG, Pereira LS, Raes D, Smith M (1998) Crop evapotranspiration: Guidelines for computing crop water requirements. FAO Irrigation and Drainage Paper No. 56, FAO, Rome 300.
21. Irmak S, Howell TA, Allen RG, Payero JO, Marti DL (2005) Standardized ASCE

- Penman-Monteith: Impact of sum-of-hourly vs. 24-hour timestep computations at reference weather station sites. *Transactions of the ASAE* 48: 1063-1077.
22. Snyder R, Pruitt W (1985) Estimating reference evapotranspiration with hourly data. Chpt VII. In Snyder R et al. (eds.) *California Irrigation Management Information System Final Report*. Land, Air and Water Resources University of California Davis, California, USA.
 23. Ortega-Farias SO, Cuenca RH, English M (1995) Hourly grass evapotranspiration in modified maritime environment. *Journal of Irrigation and Drainage Engineering* 121: 369-373.
 24. Irmak S, Payero JO, Martin DL, Irmak A, Howell TA (2006) Sensitivity analyses and sensitivity coefficients of standardized daily ASCE-Penman-Monteith equation. *Journal of Irrigation and Drainage Engineering* 132: 564-578.
 25. Tolk JA, Howell TA, Evett SR (2006) Nighttime evapotranspiration from alfalfa and cotton in a semiarid climate. *Agron J* 98: 730-736.
 26. Caird MA, Richards JH, Donovan LA (2007) Nighttime stomatal conductance and transpiration in C3 and C4 plants. *Plant Physiol* 143: 4-10.
 27. Djaman k, Irmak S, Sall M, Sow A, Kabenge I (2017b) Comparison of Sum-of-hourly and daily time step standardized ASCE Penman-Monteith (ASCE-PM) grass-reference evapotranspiration in Western Africa. *Theoretical Applied Climatology*, pp: 1-11.
 28. Bakhtiari B, Khalili A, Liaghat AAM, Khanjani J (2009) Comparison of daily with sum-of-hourly reference evapotranspiration in Kerman reference weather station, *Journal of Water and Soil* 23: 45-56.
 29. Snyder RL, Eching S (2006) PMhr Penman-Monteith Hourly ETref for short and tall canopies. University of California, Davis, California, USA.
 30. Gavilán P, Estévez J, Berengena J (2008) Comparison of Standardized Reference Evapotranspiration Equations in Southern Spain. *J Irrig Drain Eng* 134: 1-12.
 31. Suleiman AA, Hoogenboom G (2009) A comparison of ASCE and FAO-56 reference evapotranspiration for a 15-min time step in humid climate conditions. *J Hydrol* 375: 326-333.
 32. Jia X, Steele D, Hopkins D (2008) Hourly Reference Evapotranspiration Estimates for Alfalfa in North Dakota. *World Environmental and Water Resources Congress* 2008, pp: 1-10.
 33. Allen RG, Walter IA, Elliott R, Mecham B, Jensen ME, et al. (2000) Issues, requirements and challenges in selecting and specifying a standardized ET equation. *Proc 4th National Irrig Symp*, pp: 201-208.
 34. Perera K, Western A, Nawarathna B, George B (2015) Comparison of hourly and daily reference crop evapotranspiration equations across seasons and climate zones in Australia. *Agric Water Manage* 148: 84-96.
 35. Berengena J, Gavilan P (2005) Reference ET estimation in a highly advective semi- arid environment. *J Irrig Drain Eng* 131: 147-163.
 35. Berengena J, Gavilan P (2005) Reference ET estimation in a highly advective semi- arid environment. *J Irrig Drain Eng ASCE* 131: 147-163.
 36. Irmak S, Odhiambo LO, Specht JE, Djaman K (2013) Hourly and daily single and basal evapotranspiration crop coefficients as a function of growing degree days, days after emergence, leaf area index, fractional green canopy cover, and plant phenology for soybean. *Transactions of the ASABE* 56: 1785-1803.
 37. Irmak S, Djaman K, Rudnick DR (2016) Effect of full and limited irrigation amount and frequency on subsurface drip-irrigated maize evapotranspiration, yield, water use efficiency and yield response factors. *Irrig Sci* 34: 271-286.
 38. Baozhong Z, He C, Di X, Fusheng L, Show M (2017) Methods to estimate daily evapotranspiration from hourly evapotranspiration. *Bio systems Engineering* 153: 129-139.
 39. Treder W, Klamkowski K (2017) An hourly reference evapotranspiration model as a tool for estimating plant water requirements. *Infrastructure and Ecology of Rural Areas*.
 40. Colaizzi PD, Evett SR, Howell TA, Tolk JA (2006) Comparison of five models to scale daily evapotranspiration from one-time-of-day measurements. *Trans ASABE* 49: 1409-1417.
 41. Djaman K, Irmak S (2013) Actual crop evapotranspiration and alfalfa- and grass- reference crop coefficients of maize under full and limited irrigation and rain fed conditions. *Journal of Irrigation and Drainage Engineering* 139: 433-446.