

Research Article

Origin of a megascale-size elongated polje, other karst landforms and non-karst landforms in Chester County, Pennsylvania, USA

Michael lannicelli*

Brooklyn College (C.U.N.Y.), Earth and Environmental Studies, 3137 Ingersoll Hall, 2900 Bedford Avenue, Brooklyn, N.Y. 11210

Abstract

New constructive revelations arise after a controversy erupts concerning a hypothesis by Clausen [1], in which the current study proposes a counter thesis that will ultimately invalidate his hypothesis. This is because atypical landforms identified here such as an elongated polje, nested poljes and uvalas besides typical landforms, all within the Piedmont physiographic province in Chester County, Pennsylvania (PA), USA, were claimed by Clausen [1] to have originated due to a speculative, paleo-glacial ice margin and paleo-flooding. The current study considers that his hypothesis is not based on documented proof, but rather, only on unsubstantiated conjecture. Thus, the aim of the current study is to establish a feasible resolvement to the landforms' origin based on factual evidence. New revelations are ascertained here about some of these landforms being related to karst processes including influences by tectonics, structure and / or particular, geomorphological processes. The previously-mentioned, elongated polje is assigned here to Chester Valley in Chester County, PA which is a mega-scale landform encapsulating most of the landforms discussed in the current study.

Keywords: Elongated polje; Nested polje: Nested uvala: Karst; Piedmont physiographic province; Chester valley in pennsylvania

Introduction

Newly revealed, differentiated, atypical, karst landforms are recognized / identified in the current study while other typical, nonkarst landforms in the study are rightfully accorded to geomorphological processes within the Piedmont physiographic province of Pennsylvania (PA), specifically Chester County. This was prompted in the current study because a series of landforms were misinterpreted by Clausen [1] in which he claimed these were ultimately created by a speculative paleo-glacial margin and paleo-flooding. The current study looked over the geological medium that published Clausen [1] and did not see anywhere in it about if it was peer-reviewed or not peer-reviewed, but nevertheless, Clausen [1] still remains a part of the geological literature. Thus, a justifiable reinterpretation of his hypothesis will be proven in the current study while instead, a suggested, reasonable, alternative hypothesis is offered. One simple outlined strategy is to express contentions made here in the same sequence as Clausen's [1] hypothetical assertions were sequenced, so that comparisons between both published papers (his and mine) can be easily made and judged by the reader. Hence, the non-karst landforms are discussed first, followed by newly discovered, karst landforms which are then followed by a discussion about an atypical non-karst, landform. It will be proven that the set of karst landforms is identified / recognized as an elongated polje enclosing either a nested polje or a nested uvala while a relationship to the non-karst, landform assemblage will be shown in the current study. The related non-karst landforms discussed later are water gaps, wind gaps, drainage divides and barbed tributaries, all formed by stream piracy. The remaining, atypical, non-karst landform that is not related to the above landform assemblage happens to be an old but actively persisting landform that envelopes a set of numerous, SE-trending watercourses, formed during the time of the Mesozoic Period. Thus, after weighing all of the facts associated with the cumulative landforms, a credible resolution will be offered here that provides an adequate, elucidation for all of the landforms and their formative processes vs. the paleo-glacial and paleo-flooding hypothesis suggested by Clausen [1].

Background

The whole conglomeration of landforms that will be discussed in the study is pointed out in USGS topographic maps by Clausen [1]. Those landforms are all reflected by topographic contours in Clausen's [1] maps while all of these are located in Chester County, PA. The focal point of the debate takes place in this county of PA which is located far south of both the late Wisconsinan glacial border and both the Illinoian and Pre-Illinoian till limits mapped out by Sevon and Braun [2] (Figure 1). The position of Chester County, PA located within un-glaciated terrain is an integral part of the whole debate, while it is counter to Clausen's [1] rationale of his studied landforms being created either directly or indirectly by a speculative or hypothetical paleo-glacial margin that lies far south of the aforementioned glacial and till limits in PA. It is deemed necessary to first describe the geomorphology of each landscape unit, followed by arguments raised about the origin of them.

Methodology

The writer bases the current study on scientific data and facts gathered from: primary, literature sources such as many, scholarly, peer-reviewed journal papers, monographs, published geological survey reports, etc. which support the study's thesis. The cumulative data is complemented by employing bedrock, topographic, and surficial material maps that provide additional evidence and credence. Established, geomorphological models are borrowed,

*Corresponding author: Michael Iannicelli, Brooklyn College (C.U.N.Y.), Earth and Environmental Studies, 3137 Ingersoll Hall, 2900 Bedford Avenue, Brooklyn, N.Y. 11210, USA, Tel: 1-718-259-5574; E-mail: michiann@optonline.net

Received: 29-Apr-2022, Manuscript No. jescc-22-62229; Editor assigned: 02-May-2022, PreQC No. jescc-22-62229 (PQ); Reviewed: 16-May-2022, QC No. jescc-22-62229; Revised: 20-May-2022, Manuscript No. jescc-22-62229(R); Published: 27-May-2022, DOI: 10.4172/2157-7617.1000618

Citation: Iannicelli M (2022) Origin of a megascale-size elongated polje, other karst landforms and non-karst landforms in Chester County, Pennsylvania, USA. J Earth Sci Clim Change, 13: 618.

Copyright: © 2022 lannicelli M. This is an open-access article distributed under the terms of the Creative Commons Attribution License, which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.



Figure 1: Map of Pennsylvania (PA), USA showing the late Wisonsinan continental, glacial limit and also till limits deposited by Illinoian and Pre-Illinoian, continental glaciers [2]. Chester County, PA is located in the southeastern lower corner of the state Image (courtesy of DCNR, Bureau of Topographic and Geologic Survey).

correlated and applied to the study area, which happen to fittingly match up to observational landforms. The combined preceding ultimately accomplishes a skillful rationalization that leads to a valid reinterpretation of Clausen [1].

Discussion: Landform Units and Geomorphic Processes

Water and wind gaps

Water gaps and wind gaps are sizable landforms that distinctively stand out within a landscape while they've been well-known for a long time by geomorphologists [3]. Bates and Jackson [4] define a water gap as "A deep pass in a mountain ridge through which a stream flows, especially a narrow gorge or ravine cut through resistant rock by an antecedent stream". Although not specifically mentioned in the preceding definition, it is noted that this definition includes the type of stream running through the ridge, which is simply a transverse stream (as opposed to a parallel stream). A wind gap is a former water gap that does not have a stream running through it at the present time. A popular concept that explains the reason for both a present-day watercourse and only a former watercourse running through it, is headward erosion on both sides of the ridge (or a drainage divide) until a pass is lowered enough to allow a transverse stream to flow through a present-day, deepened pass. The preceding implies "stream piracy or capture" while the primary mechanism of it is when one stream's base level is lower than the base level of the captured stream at a point of intersection. Clausen [1] discussed stream piracy and headward erosion within a framework of paleo-flooding as the reason that caused these types of erosional landforms in the Piedmont province. For instance, he claimed that East Branch Brandywine Creek eroded across geological structure (which is a water gap) instead of the creek having an affinity for flowage only upon and within soft bedrock. But the current study prefers the work of Clark [3] as a wise explanation for this controversial piece. Without connotating any association with paleo-flooding, Clark [3] calculated the water gaps and wind gaps of the Piedmont Province were not only formed by transverse fluvial drainage, stream piracy and headward erosion, but that these were determined by other combinatorial factors such as tectonics and local superposition upon non-resistant cover-masses onto structural weaknesses.

Barbed tributaries, drainage divides, and negligible paleo flooding

The hypothesis of Clausen [1] attempted to explain the origin of barbed tributaries and drainage divides in Chester County, PA by correlating these to an external agent such as paleo-flooding that later initiated stream piracy. But stream piracy can result during or just after an internal agent such as a tectonic / seismic pulse that ultimately forced a degree of crustal deformation, while examples of this will be given later. Barbed drainage is a fluvial landform created from stream

piracy while the standard sequential definition of it is a stream pattern consisting of tributaries forming obtuse angles to their main stream due to these going in the opposite direction but still emptying into their main stream (Figure 2). Here, afterwards, the stream reversal ultimately pirates the adjacent stream valley when headward erosion lowered and opened up the drainage divide to the advancing, reversed stream (Figure 2) which inevitably results in a new drainage divide. When we add Clausen's [1] flooding dynamics to the equation without any tectonic implication, then sequential, hypothetical flooding of a bottomland, master, trunk River (not shown in Figure 2) causes a fluvial "push" into a tributary in the upstream direction which ultimately forces it to flow in a reverse direction. Then, after headward erosion breaches a drainage divide, the reversed stream flows through it which captures the neighboring stream. Hence, in this fashion, he implies that a newly formed drainage basin with a drainage divide is started. However, in the case of Chester County, PA, Clausen [1] did not support his own belief with any sedimentary evidence, including potential slack water deposits usually created by most flooding events as it proceeds upslope / upstream along the tributary that was eventually reversed. Slack water evidence would also record the paleo-altitude of heightened floodwaters [5, 6]. Further supporting the evidence against Clausen [1] are other investigators who don't even attribute paleo-flooding as a cause, instead, they prefer other reasons for stream piracy, stream reversal and new drainage divides such as tectonic uplift: Segar & Alexander [7] favored regional tectonic uplift for stream piracy in southeastern Tibet; Laverini et al. [8] blamed neotectonics and litho-structural patterns for both stream piracy and stream reversal in southeast Brazil; Simoes et al. [9] concluded that active tectonics are still today responsible for perpetual, drainage-divide mobility and stream capture in Bhutan, Himalaya.

Hypothetical paleo-glacier = unlikely paleo-glacier

To account for why paleo-flood sediments are absent in Clausen's [1] hypothesis, he says in a long quotation (verbatim): "Further, most melt water from such a 1 to 3 mile thick or thicker continental ice sheet would contain little or no sediment as it flowed from the ice sheet upper layers. Such melt water would deeply erode regions over which it flowed and leave little or no sediment evidence to indicate the water source. Logically massive floods of such melt water could have crossed the Schuylkill River-East Brandywine Creek drainage divide and represent the most likely erosion agent responsible for erosion events this paper describes". But Clausen's [1] rationale is grossly implausible because of a couple of reasons: Sevon et al. [10] calculated that paleocontinental glaciers were only < 0.48 km thick at Hudson Bay, Canada before moving down to Pennsylvania; while many mountain heights are > 0.48 km in the northeastern-half of the USA which would have made nunataks out of those mountains when ice flowed around them but not over them. This tenable scenario then provides a big natural chance for mobile glaciers to collect supraglacial, unconsolidated sediment due to mass movements and rock falls from frost-wedging / frost-shattering of mountainsides. In addition, atmospheric dust (loess) would have settled onto glacial surfaces [11] during wind-diminishment and dry intervals of time, while afterwards, simple snowfalls would have buried the dust during wet times. Altogether, this then contradicts Clausen [1] because supraglacial meltwater would have transported the supraglacial sediment from off the surficial part of the glacier while fluvioglacially depositing at least some of it onto the terrain, especially in slackwaterdeposit positions within a paleo-flooding scenario ----- which wasn't the case here.



Figure 2: Two-stage diagram showing the development of stream piracy and barbed tributaries. In diagram-1, headward erosion downcuts upon both sides of a drainage divide, while downcutting also proceeds downslope within two separate, dendritic, drainage basins on each side of the drainage divide. In diagram-2, active tectonics force stream reversal of the main stream that occurs in at least one of the drainage basins. This incurs obtuse angles to the tributaries relative to the main stream. The reversed main stream now flows through the drainage divide that was eroded down earlier by headward erosion, ultimately pirating the adjacent stream. Both stream piracy and barbed drainage if caused by a hypothetical flooding event is explained in the text of the current study. Image courtesy of: https://www.geographynotes.com/drainage-system/10-main-types-of-drainage-patterns-streams-geography/2425.

J Earth Sci Clim Change, an open access journal ISSN: 2157-7617

An elongated polje and nested poljes and uvalas

Clausen [1] says two streams run parallel to one another within the NE-SW- trending, megascale-size, lowland of Chester Valley in PA [12] (Figure 3), while each one flows in opposite directions, one (Valley Creek West) towards the W and the other (Valley Creek East) towards the E both separated by a divide (Figure 4). The landform "throughvalley" was designated for these valleys in Clausen [1] but he did not differentiate it from two different types of landforms that are both known as a "through-valley". Thus, the current study analyzes both landform types to cover both possible versions of the through-valleys in the study area. Clausen [1] claimed that a hypothetical paleo-glacial margin operated closely in the study area, which motivates the current study to explain the sub-topic of "glacial through-valleys". This variety of through-valley evolves when glacial ice exploits a divide enough to soften gaps within the divide which results in uniting opposite, oriented streams into one continuous valley as reported by Coates [13]. But one good reason against Clausen's [1] speculation of paleo-glacial ice as being the erosional agent responsible for his through-valley is simply the location of his study area at a very far distance away from the late Wisconsinan glacial limit and the till limit deposited by Illinoian and pre-Illinoian glacial ice (Figure 1) thus, both true glacial-ice limits and till limits were not respected by Clausen [1]. The other type of "through-valley" which will now be discussed is correlated to karst or carbonate-rock topography. So, to recapitulate, Clausen [1] discussed the two, parallel creeks that are separated by an N - S divide while flowing in opposite directions which is Valley Creek West and Valley Creek East. Each creek debouches onto Chester Valley via water gaps (Figure 4) that transversely cut through the south side, valley wall of Chester Valley underlain by phyllitic-shale of the Octararo Formation striking generally ENE - WSW. But Clausen [1] believed at one time in the past, both creeks flowed in the same direction which convinced him that one of the two streams had to be reversed during the past due to paleo-flooding. Stream reversal and its mechanism of origin have already been discussed and concluded in one of the earlier sections of the current study, so this is applied to here as well, which results in eliminating his belief on this particular matter.

A feasible origin of Valley Creek West and Valley Creek East (Figure 4) is ascertained here by first determining an origin for the megascalesize Chester Valley (Figure 3) as a whole landscape unit because it may shed light on those two creeks. Kochanov [14] did not offer a reason for the creation of Chester Valley in PA., but Bascom et al. [15] did, since he said that a river never ran through it, but instead, chemical weathering caused dissolution of the carbonate bedrock underlying Chester Valley which resulted in its geomorphic form. The current study agrees with Bascom et al. [15] and now analyzes this megascalesize landform to see if we can deduce a true, complete picture of its origin. First, a description needs to be given about the underlying structure and strata of the study area in Chester Valley. The structural foundation of it is a leveled series of asymmetric and symmetric synclines and anticlines trending transversely across the valley with the inclusion of various types of faults which altogether give the underlying lithostratigraphy of Chester Valley, a slightly complex, deformational pattern. The age of the faults was not determined by Kochanov [14]. The surficial rock lithologies immediately underlying Chester Valley are either one of two carbonate, Cambrian-age, dolostone units which are the Ledger Formation and the Elkbrook Formation plus



Figure 3: Chester County Rock-Type Map in Pennsylvania [12]. Chester Valley is the WSW – ENE elongated bedrock lowland that stretches across the whole map, delineated by two hues of purple symbolizing different bedrock formations of dolomite and limestone. A topographic look at one part of the lowland is in Figure 4. There are different types of faults underlying Chester Valley which are not seen on the rock-type map, but refer to Kochanov [14, his Plate 1] to see these, as well as the associated, structural folding of bedrock formations Image courtesy of Pennsylvania Department of Conservation and Natural Resources.

J Earth Sci Clim Change, an open access journal ISSN: 2157-7617



Figure 4: Topography of allogenic creeks Valley Creek West and Valley Creek East flowing towards the N through and from out of their own water gaps (at the bottom of the map) which cut across a phyllitic-shale, bedrock ridge. They are partitioned by an N-S bedrock divide (South Valley Hills) which then extends and continues N and transversely over the vast carbonate lowland of Chester Valley. As each creek flows over the carbonate lowland, Valley Creek West makes an abrupt turn towards the W while Valley Creek East makes its own abrupt turn where it intersects Swedesford Road and towards the NE. Parallelism is always maintained between the two creeks. See text for additional details. From the Malvern, PA 7.5' topographic quadrangle map, courtesy of the USGS.

in addition, Ordovician-age limestone bedrock of the Conestoga Formation. These three stratums are overlain by Quaternary sediments of colluvium, alluvium, residuum and saprolite. As mentioned earlier, the lowland of Chester Valley's carbonate bedrock floor is in contact with its own valley walls that are composed of phyllitic shale along its southern side. The lengthy, phyllitic-shale, valley wall was punctuated and completely eroded through by many allogenic creeks, leaving behind a very long, ENE - WSW trending string of ubiquitous, N - S oriented, water gaps (Figure 4). This is while Chester Valley's northern side is walled with an unnamed unit - mix of granodiorite and gneisses, plus quartzite of the Chickies Formation. The size of Chester Valley is generally 6 km in width and 88 km in length and it is noted here that the megascale size of it extends beyond Chester County, PA and into neighboring Counties of PA. The current study interprets Chester Valley with its combinatorial characteristics as a "polje". Selby [16] described two different types of poljes: one type is a very large (in areal extent), flat-floored, depression in carbonate bedrock bounded all around by a vertical-walled scarp (Figure 5) due to either half-grabens and full grabens that initiated by either gravity faulting or tectonic faulting; another type is also very large in size, having its carbonate bedrock underlain by either a syncline or anticline (Figure 5). Besides the structural influences upon the carbonates involved in the geomorphic development of poljes, other factors are considered as well, such as differential erosional rates between the contact of different bedrock formations through dissolution and both surface drainage and subsurface drainage, as reported by Milanovic [17].

Poljes are intermontane plains that are well-known in the Balkan countries and in the Middle East such as Turkey, while their sizes can be up to 700 km² in area [18]. Dimension-wise, the largest poljes in the world were mapped out by Selby [16] who illustrated lengths of

70 km within the Dinaric Mountains of the Balkan Peninsula. Some poljes are water-filled or flooded due to ponors within their carbonate floors being choked with debris such as alluvium. Traditionally, the general perception of polies is that they are circular or oval in shape, but in fact, a polje may also be only of an elongated shape or something intermediate in form between a circular and elongated shape [19] (Figure 6). Thornbury [20] besides Selby [16] said poljes do not evolve from karstic "uvala" landforms, but rather, due to structural erosion upon a large-scale anticline / syncline or upon faulted blocks where there is contact between insoluble bedrock and soluble bedrock, as mentioned earlier. Uvalas are either relatively short or long bedrock depressions that develop simply through dissolution due to a collapse of many, curving, inter-doline ridges belonging to a string of dolines without the influence of any structural control. It is tempting to say here that Figure 6 represents evolutionary phases of an elongated polje transitioning into a circular polje but this may not be the case because Simsek and Garcia et al. [21] report elongated poljes are dictated by parallel, tectonic, structural control while circular poljes are influenced by fluviokarst. However, it is not certain here how the latter mechanism works in that regard. Anyway, the point of this particular discussion is that the characteristics and the present shape of Chester Valley in PA meet the same criteria as an elongated-shaped polje.

Selby's [16] illustrated model of his first, polje type (Figure 5) is very similar to the paleogeomorphic setting of Chester Valley because of insoluble bedrock (the shale) uplifted along the contact with the polje's carbonate floor while allogenic streams head from off the uplifted highland and onto the lowland of the polje. In Chester Valley's case, headward erosion by allogenic creeks eventually eroded vertically and completely downward upon the paleogeomorphic, phyllitic-shale, and valley's wall in spots which resulted in today's Valley Creek West

Page 5 of 8

Page 6 of 8



Figure 5: Particular karst depicted by two different types of poljes. Each polje is underlain by carbonate bedrock while structurally controlled by either faulting seen in the top diagram or a syncline / anticline seen in the lower diagram. Erosional, dissolutional processes combine with fluvial processes in conjunction with the structural controls to create a polje. Thus, poljes are differentiated from other typical karst landforms because the latter is formed strictly from only dissolutional processes of carbonate bedrock from Selby [16].



Figure 6: Plan view of various-shaped poljes located in the Taurus Mountains of southern Turkey [19]. The elongated shape of Chester Valley in Chester County, PA (Figure 3), is akin to either of the several, elongated poljes on the left-half side of the diagram. Names of the poljes from left to right are: Kovada Polje, Karadiken Polje, Gembos Polje, Camell Polje, Geyron Polje, Karakose Polje, and Akyaka Polje. Scale: length of Gembos Polje is ca.12 km long while the diameter of Akyaka Polje is < 12 km from: Simsek [19].

and Valley Creek East cutting initially and transversely across Chester Valley (Figure 4). Hence, it should be stressed again that Figure 5 provides an excellent glimpse into the paleogeomorphology of Chester Valley before the creation of its water gaps. The carbonate lowland of Chester Valley is largely in an inactive phase today by evidence of relict, karst features in it such as: the Port Kennedy Bone Cave (which is actually a large sinkhole of at least Pleistocene age); the surficial dolostone of the Ledger Formation which is a sinkhole-prone, highly-soluble, highly fractured, bedrock; and the inclusion of of .5 m-wide ponors. Thus, it's reasonable to presume that the paleoclimate in the region was more humid than today when it generated those erosional, carbonate-rock features.

Before we round out an accurate origin for the previouslydiscussed, confounded, "through-valley" containing Valley Creek West and Valley Creek East within the mega-scale size Chester Valley, an additional discussion is imperative. The through-valley is perceived in the current study as one or both of the following karst landforms which are a "nested polje and uvala". A nested polje is defined by Sauro [21] as simply a small polje enclosed by a large polje as illustrated by him (Figure 7). The smaller landform there is indeed nested, but it is doubtful about it being a smaller polje because there is no structural influence associated with this nested landform seen in the cross-section of Figure 7. Instead, it is most probably an uvala that is erosionally superimposed over the polje simply through dissolutional processes of the limestone bedrock in Sauro's [21] illustration (Figure 7). A detailed description presented here now of the topography associated with Valley Creek West and Valley Creek East may help us affirm a minutiae of their genesis. As seen in Figure 4, each creek flows through their own water gap that is partitioned by an N-S divide (labeled as "South Valley Hills" on the map) extending from the highland of the valley wall and continuing onto the vast lowland of Chester Valley. Here, Valley Creek West enters the lowland in an N direction but then abruptly curves toward the W and runs down a gentle slope. In slight contrast, Valley Creek East enters the lowland in a N direction but then flows over a semi-circular, one-sided collapse-depression (karst?) that seemingly causes it to abruptly curve towards the E, followed by flowage over a slender-shaped, collapse-depression (karst?) and then down a

J Earth Sci Clim Change, an open access journal ISSN: 2157-7617



Figure 7: Cross-sectional view of a large, inactive polje with a nested karst landform that was eroded into a part of its surface [21]. It is located within the Piano del Cansiglio

Figure 7: Cross-sectional view of a large, inactive polje with a nested karst landform that was eroded into a part of its surface [21]. It is located within the Piano del Cansiglio in the Venetian Prealps of Italy. Note the analogy of the inactive polje to Chester Valley in Chester County, PA because of similar carbonate bedrock lithologies, folded strata, and faults, plus its "nested polje" is geomorphically similar to either Valley Creek West and Valley Creek East that are both enclosed by Chester Valley. Permission to reproduce is granted by Elsevier Science, License No. 5231511216467.

declivitous slope. The study interprets the characteristics of both creeks may be related to stream flowage guided over the previously-mentioned, series of asymmetric and symmetric synclines attached to a chain of synclines and anticlines that transversely underlie Chester Valley. If this is truly the expected association, then Valley Creek West and Valley Creek East are identified as "nested poljes" situated within the megascale-size, elongated polje which is Chester Valley. Alternatively, if we take into account about the existence of two collapse-depressions underlying small parts of Valley Creek East, then we cannot rule out an uvala origin, which actually means that a "nested uvala" is erosionally overprinting a portion of the much bigger, elongated-shaped polje that constitutes Chester Valley.

The literature about poljes in the USA

The literature's identification of poljes in the USA is somewhat muddled. According to Milanovic [17], he was told by a prominent geomorphologist (Richard Parizek) that poljes exist in both Centre County and Lycoming County of Pennsylvania while the largest of these is Phantom Lake. Much farther south, Klindinger and Flocks [23] said there is a peculiar, irregular-shaped, active polje named Orange Lake, in north Florida which sits on the Florida Upland physiographic region flanked on both sides of it by the Coastal Plain. They said it formed through: dissolutional erosion of Florida's Tertiary-age, carbonate platform caused by a deep, 50 m-wide, collapsed sinkhole (with small-scale faulting) and an adjacent, surrounding, subsidence sinkhole that forced a shallower de-elevation of the surficial, limestone bedrock containing thick vertical, solution pipes within bedrock fractures, altogether without any associated, structural folding. But in the case of Orange Lake, the term "polje" may be a misnomer in Klindinger and Flocks [23] because even though it has down-faulted limestone, it lacks any contact with insoluble rock nor does it possess any folded bedrock which altogether contradicts a polje definition for it. As a side note, they also said that Orange Lake is colloquially known only as a "drowned prairie". Closer to the strict definition of "poljes", are the oval-shaped type of poljes in the USA which are assigned here to the "coves" (and not "piedmont coves", sensu stricto, in the Blue Ridge province, see Mills [24], his figure 2) such as "Cades Cove" [25] within the Great Smoky Mountains of Tennessee. This is because besides it having nearly the same circulat to oval, geomorphic shape as a polje in Turkey, it also possesses the same properties because of subsurface folds, and different types of faults including similar lithological bedrock units. So, Cades Cove is comparable to Chester Valley in PA, but there are a few differences here which are: the former has an oval shape while being located in the Blue Ridge province vs. the latter which has an elongated shape while being located in the Piedmont province; and dimensions between the two are significantly different. It should be noted that in the adjacent, structural province to the Blue Ridge, which is the Ridge and Valley province, there is the "Dungannon Polje", as only mentioned by Clark et al [26]. Also, it's worth mentioning here about a similar landform known as a tectonic window or fenster, but here, the standard definition of it does not necessarily include any association to soluble bedrock lithologies.

Page 7 of 8

In summing up here, we can rule out paleo-glacial ice of any age in creating the E - W, narrow and shallow valleys of both Valley Creek West and Valley Creek East, while designating these as either "nested poljes", or in the case of Valley Creek East, only as a "nested uvala" which is another appropriate possibility. The erosional landforms underlying both creeks are encompassed by the whole landscape unit called Chester Valley while it is formally identified here in the current study as an elongated polje.

SE-trending watercourses

Clausen's [1] hypothesis includes "valley orientations" (termed by him), for example, the following SE-trending, rivers, streams, creeks and tributaries: the Schuylkill River, East Branch Brandywine Creek, West Branch Brandywine Creek, Wissahickon Creek (north of Philadelphia, PA) and several Octoraro Creek tributaries, all within Chester County, PA. He believed these resulted from SWflowing, floodwaters originating from a fast-melting, paleo-glacial margin belonging to a speculative, ancient, continental ice sheet that operated far beyond (south of) the late Wisconsinan glacial margin and even far beyond (south of) the till limit of Illinoian and Pre-Illinoian continental glaciers (Figure 1). But to the contrary, Clark [3] explains the origin of the SE-trending watercourses by referring to Faill [27] who documented sedimentological evidence supporting a Mesozoic-age, sub-aerial, deltaic, fan which hosted and influenced the previously-mentioned, SE-trending Schuylkill River within the structural Newark-Gettysburg Basin. Clark [3] also included a major trunk valley containing the SE-oriented Susquehanna River (located outside of Chester County and within PA) as flowing on the paleo-deltaic fan too during Mesozoic time. These two rivers do show up on the map of Sevon and Braun [2] although they are not labeled on this map. Thus, both SE-trending watercourses are correlated to all of the above previouslymentioned, modern-day, SE-trending watercourses which inherited their original SE direction from paleo-fluvial flowage upon a deltaic fan that survived through the ages of geologic time, while these simply kept actively flowing within entrenched channels in the same SE direction up to the present time.

Conclusions

The current study gives ample proof negating Clausen's [1] hypothesis of a paleo-glacial margin and paleo-flooding in Chester County, PA, USA as an origin for the discussed assemblage of different, erosional landforms. It is demonstrated here that other grouped, geological factors were responsible for most of the landforms' origin such as a combination of tectonics, structure, lithology, and basic geomorphological processes. Even more important, the current study uncovered new revelations by identifying the very large, megascalesize, Chester Valley in PA as an elongated polje. Associated with that, the true nature of the "through-valley" (termed by Clausen [1]) is identified as either a "nested polje" or a "nested uvala". Nested poljes most probably embody both Valley Creek East and Valley Creek West because of a series of small, leveled, symmetric and asymmetric synclines underlying Chester Valley even while Chester Valley's origin is structurally controlled by only a fault, rather than to the underlying synforms associated with the creeks. Meanwhile, an alternate explanation is considered for only Valley Creek East because it may be a "nested uvala" since the possibility of only paleo-dissolutional processes overprinted a part of Chester Valley's polje. It's emphasized here that the key to comprehending the genesis of Chester Valley was originally hinted at by Bascom et al. [15] because they were the first investigators to say the megascale-size valley was not formed by a river running through it, but instead, formed by wholesale, carbonate dissolution of it. Overall, both the discussed karst landforms and non-karst landforms of the current study are all related to tectonics, structure and particular geomorphological processes. Only one of the discussed landforms in the current study do not share the same, above, geological characteristics which is the grouped, SE-trending watercourses that have a remarkable, inherited, persisting quality to them, ongoing ever since Mesozoic times. Clausen [1] was honest enough to admit having many unanswered questions about his own hypothesis, in which case, supplements the nullification given here, as verified by the current study.

Acknowledgements

A suggestion made by an anonymous reviewer at the Journal of Cave and Karst Studies was ultimately incorporated within the manuscript of the author's study. The funding source was completely sponsored by the author of this article.

References

- Clausen E (2017) Origin of erosional landforms along and near the Schuylkill River-East Branch Brandywine Creek drainage divide segment of the Schuylkill River-Delaware River drainage divide, Chester County, PA. In: Bosbyshell, H. (Ed.), GANJ (Geological Association of New Jersey), XXXIV. Annual Meeting and Field Guide.
- Sevon WD, Braun DD (1987) Glacial deposits of Pennsylvania. Pennsylvania Geological Survey, Map 59.
- 3. Clark GM (1989) Central and southern Appalachian water and wind gap origins: review and new data. Geomorphology 2: 209-232.
- Bates RL, Jackson JA (1987) Glossary of Geology. 3rd Edition, American Geological Institute, Alexandria. 778 p.
- Ritter DF, Kochel RC, Miller JR (1995) Process geomorphology: Dubuque, IA: W.C. Brown Publishers, 539 p.

- Anderson RS, Anderson SP (2010) Geomorphology: the mechanics and chemistry of landscapes. Cam Uni Press, UK, New York and other locations, 637 p.
- Seger M, Alexander J (1993) Distribution of Plio-Pleistocene and Modern coarse-grained deltas south of the Gulf of Corinth, Greece. Intl Assoc of Sedimentologists, Blackwell Sci Pub, UK and elsewhere p. 37-48.
- Laverini C, Magalhaes Jr. AP, de Oliveira FS, de Carvalho A (2002) Neotectonics, river capture and landscape evolution in the highlands of SE Brazil. Mercato Fortaleza 15(4): 95-119.
- Simoes M, Sassolas ST, Cattin R, Leroux MR., Ferry M, et al. (2021) Topographic disequilibrium, landscape dynamics, and active tectonics: an example from the Bhutan Himalaya. Earth Surf Dyn 9; 895-921.
- Sevon WD, Fleeger GM, Shepp VC (1993) Pennsylvania and the Ice Age. Pennsylvania Geological Survey, Educational Series 6, 30 p.
- 11. Benn DI, Evans DJA (1998) Glaciers and glaciation. Hodder Arnold Pub, London, 734 p.
- PDCNR [Pennsylvania Department of Conservation and Natural Resourses] (2017) Chester County Rock-Type map in Pennsylvania, scale: 1: 36,000.
- Coates DR. (1974) Reappraisal of the glaciated Appalachian Plateau. In, Coates, D.R., (Ed.), Glacial Geomorphology. George Allen and Unwin, UK and Massachusetts, p. 205 - 243.
- Kochanov WE (2016) Geology of part of the Chester Valley area, Chester, Delaware, Montgomery, and Philadelphia Counties, Pennsylvania. Pennsylvania Geological Series, 4th Series, Open File Geologic Atlas 16-01.0, 40 p.
- Bascom F, Clark WB, Darton, NH, Knapp GN, Kuemmel HB, et al. (1909) Philadelphia folio: Norristown, Germantown, Chester, and Philadelphia, Pennsylvania-New Jersey-Delaware. Folios of the Geologic Atlas 162.
- Selby MJ (1985) Earth's Changing Surface: An Introduction to Geomorphology. Oxford, UK and New York: Clarendon Press 480 p.
- 17. Milanovic PT (2004) Water resources engineering in karst. CRC Press 330 p.
- Gutierrez F, Gutierrez M (2016) Landforms of the Earth: an illustrated guide. Springer Intl Pub, 270 p.
- Simsek M, Ozturk MZ, Dogun U, Utlu M (2021) Morphometric properties of poljes in the Taurus Mountains of southern Turkey. J Geography 42; 101-119.
- Thornbury WD (1954) Principles of geomorphology. John Wiley & Sons, Inc, New York, 618 p.
- Sauro U (2012) Closed depressions in karst areas. In: White WB, Culver DC (Eds.), Encyclopedia of caves, (2nd ed): Elsevier Academic Press, London and elsewhere, p. 140 – 155..
- Garcia FJ, Gutierrez F, Gutierrez M (2003) The Jiloca karst polje-tectonic graben (Iberian Range, NE Spain). Geomorphology 52; 215-231.
- Kindinger JL, Flocks JG (2000) Geologic controls on the formation of Florida sinkhole lakes. USGS Open-File Report 00-294, 4 p.
- 24. Mills HH (2000) The relationship of slope angle to regolith clast size: a study based on surficial mapping in the southern Blue Ridge province, western North Carolina. In, Mills, HH (Ed.), Regolith in the Central and Southern Appalachians: Southeastern Geology, 39(3 and 4): 243-258.
- Moore HL (1988) The roadside geology of the Great Smoky Mountains. The University of Tennessee Press, Knoxville, Tennessee, 192 p.
- 26. Clark GM, Ciolkosz EJ, Kite JS, Lietzke DA (1989) Central and Southern Appalachian Geomorphology. Tennessee, Virginia, and West Virginia: Maryville, Tennessee to Washington, D.C. July 2–9, 1989. AGU 105 p.
- Faill RT (1973) Tectonic Development of the Triassic Newark-Gettysburg Basin in Pennsylvania. Geol Soc Am Bull 84 (3): 725-740.